## Q1: Hurricane Physics

A hurricane is a strong tropical storm having a low-pressure center (the eye of the hurricane) around which the air masses rotate with high speed. The general name for this type of extreme weather condition is that of cyclone. In the Western hemisphere it is called hurricane, while in Asia it is called typhoon. Regardless of the name, the hurricane forms above the warm sea water.

## Task 1: Hurricane Mechanics ( $\mathbf{4 . 4 0}$ points)

To understand more about the mechanics of air movement towards the hurricane's eye, let us consider an idealized vortex - the cylindrical one. For simplicity, consider that the eye wall is cylindrical and vertical, taking its symmetry axis as the axis Oz of a cylindrical coordinate system (having $\hat{r}, \hat{\varphi}$ and $\hat{z}$ as unit vectors). The air flows horizontally, towards the eye wall, so the velocity of any fluid particle outside the eye wall is $\vec{v}=\hat{r} v_{r}+\hat{\varphi} v_{\varphi}+\hat{z} 0$ (see the adjacent figure).
Note: In Fluid Mechanics, a fluid particle is a fluid volume, much smaller than the entire fluid volume to be considered as a particle and much bigger than a fluid molecule to
 encompass all the macroscopic properties of the fluid.

| 1.a. | Determine a mathematical expression for the velocity of a fluid particle at a <br> distance $r$ form the axis, in the outer region of the vortex. The fluid particle enters <br> the outer region at a distance $r_{0}$ from the axis, with the radial speed $v_{0}$. In the outer <br> region of the vortex the air flow is not rotational. | $\mathbf{0 . 6 0}$ <br> $\mathbf{p}$ |
| :---: | :--- | :--- |

Between the outer region and the eyewall there is an annulus (a cylindrical layer) in which the velocity of the fluid particle has an azimuthal component $\left(v_{\varphi}=\frac{C}{r}\right)$, also keeping the same radial dependence of $v_{r}$, as in the outer region. $C$ is a positive, known constant.

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1.b. of the fluid particle from the outer region to the annulus takes place at
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1.70
p

Usually, a difference in temperature in the atmosphere will produce a pressure difference, and the pressure force will move the atmospheric air, creating what we commonly call winds. We witness this air circulation from the Earth, which is a non-inertial reference frame due to its diurnal rotation (with an angular speed $\omega=7,29 \cdot 10^{-5} \mathrm{rad} / \mathrm{s}$ ). In such a reference frame, to formally write Newton's second law of motion, some inertia forces should be added to the ones resulting from interactions. Among them, the Coriolis force is of interest here. Its expression is $\vec{F}_{\text {Cor }}=-2 m \vec{\omega} \times \vec{v}_{r e l}$, where $\vec{v}_{r e l}$ is the relative velocity of a body with respect to the non-inertial reference frame. Although the Earth's angular speed is rather small, the value of the Coriolis force becomes noticeable as the speed of the body increases.
In the adjacent sketch you can see a crude representation of a low pression region, some isobars (the dashed curves) and a few arrows representing the wind directions. The representation does not consider the Earth rotation around its own axis. Imagine that this sketch describes a real situation somewhere in Romania (for simplicity, consider the North latitude as $\lambda=45^{\circ}$ )

1.c. Redraw the above sketch and on each arrow draw the direction of Coriolis force;
0.50 as a result, redraw the air currents (arrows) affected by the Coriolis force. p

When the wind speed increases, its direction is more and more modified by the Coriolis force, until the fluid particles spiral around a closed loop where the Coriolis force has values comparable with the pressure force. This type of circular flow is observed in the annulus, close to the eyewall of the hurricane.
Consider a fluid particle with a small volume $\delta V$ at a distance $r$ from the eye axis, moving around it (the air density is $\rho=1.29 \mathrm{~kg} / \mathrm{m}^{3}$ ) with constant speed.

| 1.d. | Derive the equation of motion for the fluid particle, neglecting the viscosity of air, <br> and obtain a mathematical expression for the wind speed as a function of the <br> pressure gradient $\frac{\delta p}{\delta r}$ <br> the angular speed of the Earth $\omega$. | $\mathbf{1 . 3 0}$ |
| :---: | :--- | :--- |

The table below gives the so-called Saffir-Simpson intensity scale for hurricanes.

| Category | Minimum central sea- <br> level pressure/kPa | Maximum wind <br> speed/(m/s) | Damage |
| :---: | :---: | :---: | :---: |
| 1 | $\geq 98.0$ | $33-42$ | minimal |
| 2 | $97.9-96.5$ | $43-49$ | moderate |
| 3 | $96.4-94.5$ | $50-59$ | extensive |
| 4 | $94.4-92.0$ | $60-69$ | extreme |
| 5 | $<92.0$ | $\geq 70$ | catastrophic |


| 1.e. | $\begin{array}{l}\text { Calculate the minimal value of the air pressure gradient at } 25 \mathrm{~km} \text { from the center } \\ \text { of the low-pressure region to qualify a wind as a hurricane. }\end{array}$ | $\begin{array}{c}\mathbf{0 . 3 0} \\ \mathbf{p}\end{array}$ |
| :---: | :--- | :---: |

## Task 2: Hurricane Thermodynamics ( $\mathbf{5 . 6 0}$ points)

A hurricane path is not very difficult to forecast, but until recently, its intensity was. A new model, which is thermodynamical in nature, is able to offer accurate results ${ }^{1}$. In essence, a hurricane operates as a heat engine, transforming a part of the heat absorbed at the ocean surface to mechanical energy of the hurricane wind. The succession of steps can be, briefly, outlined as follows:

1. As the air currents (wind) spiral towards the eye of the hurricane, they absorb heat from the ocean's surface (water from the surface of the ocean vaporizes and is carried by the wind together with its latent heat of evaporation); the
 segment $\mathrm{A}-\mathrm{B}$ is essentially an isothermal process at the temperature $T_{1} \cong 300 \mathrm{~K}$.
2. Reaching the eye wall, the moist air will rise quickly up to about 17 km (segment $\mathrm{B}-\mathrm{C}$ ).
3. A short, isothermal segment $(\mathrm{C}-\mathrm{D})$ during which the wind transfers heat to the outer space and the air dries up (the temperature is $T_{2} \cong 200 \mathrm{~K}$ ).
4. The dry air quickly descends towards the ocean surface (segment $\mathrm{D}-\mathrm{A}$ ).
[^0]| 2.a. | Using the above information, determine the maximum rate of energy transferred <br> from the ocean to the kinetic energy of the wind, considering that the absorbed <br> heat rate from the ocean surface is $\frac{d Q_{1}}{d t}>0$. | $\mathbf{0 . 8 0}$ <br> $\mathbf{p}$ |
| :---: | :--- | :---: |

In a hurricane, the mechanical energy of the wind is converted into heat due to wind friction, almost all of it at the ocean surface. This heat, in turn, contributes to the rate at which heat is absorbed by the wind at the ocean surface.
The entire air circulation takes place inside a cylindrical volume, having the hurricane eye axis as its symmetry axis, the height of 17 km , and the radius of its circular base $R=100 \mathrm{~km}$.

| 2.b. | If $d Q_{10} / d t$ is the rate at which the heat is released from the ocean's surface by <br> evaporation, write an expression for $\frac{d Q_{10}}{d t}$, relating it to the given temperatures, <br> the air density, and the wind speed. Any non-dimensional proportionality <br> coefficient being of the order of unity, should be taken as equal to one. | $\mathbf{1 . 6 0}$ |
| :---: | :--- | :---: |
| $\mathbf{p}$ |  |  |

Let us denote by $h$ the specific enthalpy (the enthalpy per mass unit) of the inflowing dry air (at constant pressure, the enthalpy variation is the heat exchanged by the system with its environment) and with $h^{*}$ the specific enthalpy of the wet air immediately above the ocean surface (the latent heat of evaporation is, in fact, the evaporation enthalpy).

| 2.c. | Derive a mathematical expression for $\frac{d Q_{10}}{d t}$, relating it to the given specific <br> enthalpies. Any non-dimensional proportionality coefficient being of the order <br> of unity, should be taken as equal to one. | $\mathbf{1 . 2 0}$ <br> $\mathbf{p}$ |
| :---: | :--- | :--- |

Under stationary conditions (constant local pressure and temperature in any given point on the surface of the ocean's water), the specific enthalpy of water vapors at $0^{\circ} \mathrm{C}$ is $h_{w 0}=2437 \mathrm{~kJ} / \mathrm{kg}$, and the specific heat of water is $c_{p, w}=4.18 \frac{\mathrm{~kJ}}{\mathrm{kgK}}$. The air above the water surface is saturated with water vapors. The molar mass of air is $M_{a}=28.97 \mathrm{~g} / \mathrm{mol}$, saturated vapor pressure of water at 300 K is $p_{s}=3.561 \mathrm{kPa}$ and the atmospheric pressure is 101.3 kPa . The water vapors can be considered an ideal gas.

| 2.d. | Assuming that both expressions derived for $\frac{d Q_{10}}{d t}$ are valid for any value of the <br> radius $R$ greater than the radius of the eye, derive an expression for the maximum <br> wind speed and calculate its numerical value. | $\mathbf{2 . 0 0}$ <br> $\mathbf{p}$ |
| :---: | :--- | :---: |

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[^0]:    ${ }^{1}$ The author of the thermodynamic model is dr. Kerry Emanuel, from MIT.

